# Combined use of SDMT-CPTU results for site characterization and liquefaction analysis of canal levees

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ABSTRACT: The paper illustrates the combined use of the results of seismic dilatometer tests (SDMT) and piezocone tests (CPTU), obtained as part of a comprehensive study aimed at investigating the post-earthquake stability conditions of canal levees damaged by the May 2012 Emilia (Italy) seismic sequence. The following issues are discussed: (1) Comparison of results obtained by SDMT vs. CPTU test interpretation, in particular soil type identification and ground property characterization of the embankment and the foundation soils. (2) Liquefaction analysis using a recent simplified method (Marchetti 2016) based on the combined use of the horizontal stress index  $K_D$  provided by SDMT and the cone penetration resistance  $q_t$  provided by CPTU. The results obtained by this method are compared with the results obtained by existing methods based on  $q_t$  (CPT) and  $K_D$  (DMT) alone, as well as with the results obtained by methods based on the shear wave velocity  $V_S$  and by laboratory cyclic tests.

# 1 INTRODUCTION

The seismic sequence which in May 2012 struck a wide area of the Po river plain (Emilia-Romagna region, Northern Italy) caused extensive damage to a number of riverbanks in the epicentral area, in the form of ground deformations, surface fractures and lateral spreading. Major damage was observed in a 3 km long segment of the embankment bordering an irrigation canal known as "Canale Diversivo di Burana" near Scortichino, Bondeno (Ferrara), hosting more than one hundred houses and productive activities. In some cases buildings and facilities built on the bank crown were found unstable or unsafe and thus declared unfit for use.

The municipality of Bondeno, supported by the Emilia-Romagna regional authority in cooperation with the Italian Geotechnical Society (AGI), promoted a Working Group of researchers from various Italian universities and experts of the Geological, Seismic and Soil Survey Regional Department, committed to analyzing the seismic response of the embankment, investigating the causes of the earthquake-induced damage, assessing the postearthquake stability conditions and finally proposing remedial measures. A comprehensive site investigation program, including several in situ and laboratory tests, was performed for this task. The most significant results achieved by the Working Group activity

were summarized by Tonni et al. (2015a). This paper is focused on the combined use of results from seismic dilatometer tests (SDMT) and piezocone tests (CPTU) for site characterization and liquefaction analysis.

## 2 SDMT-CPTU TESTS IN THE SCORTICHINO EMBANKMENT AREA

# 2.1 Testing program and location

The canal levee and the foundation soils were extensively investigated by in situ tests (5 boreholes, 12 CPTUs, 4 SDMTs, piezometer measurements, permeability tests) and by a large number of laboratory tests (triaxial, shearbox, resonant column and cyclic torsional shear, cyclic simple shear, double specimen direct simple shear) on undisturbed and reconstituted samples. Details on the test results and the relevant geotechnical parameters can be found in Tonni et al. (2015a,b). The site investigations were concentrated along cross-sections in four distinct areas (A, B, C, D, Fig. 1), located at about 1 km distance from each other, in which the most severe and extensive damages had been observed, particularly in the area C. In each investigated area (Fig. 2) at least three CPTU/SDMT soundings were carried out from the crest of the embankment, down to 30-35 m depth.



Figure 1. Aerial view of the damaged bank stretch and location of the investigated areas.



Figure 2. Location of in situ tests in the four selected areas.

## 2.2 Interpretation of CPTU and SDMT results

The borehole logs and the interpretation of CPTU and SDMT results (Tonni et al. 2015a,b) consistently allowed recognizing the following stratigraphic sequence and soil units from the crest of the embankment (see e.g. cross section c-c', area C, Fig. 3):

- an upper soil layer, about 9-10 m thick, composed of sandy silts and silty sands, corresponding to the core of the man-made embankment (Unit AR) in the topmost 6-7 m and to natural soils (Unit B) in the bottom portion;
- a clayey-silt layer with inclusions of peat and organic material (Unit C), generally  $\approx$  1-2 m thick;
- a medium to coarse or very coarse sand layer (Unit A) extending down to the maximum investigated depth, at least 40 m in thickness, locally including thin clayey lenses at depths between 30 and 34 m from the crest of the levee.

The above soil sequence was encountered in all the investigated areas, with minor variations in thickness of distinct soil units and/or in composition (predominantly sandy or silty) of Units AR and B. As an example, Fig. 4 shows the profiles of the corrected cone resistance  $q_t$ , the sleeve friction  $f_s$  and the pore pressure u provided by the piezocone test CPTU 6, carried out from the bank crest in area C. The plot also includes results from CPTU data interpretation, i.e. the profile of the Soil Behaviour Type (SBT), based on the Soil Behaviour Type Index  $I_{cn}$  calculated from the normalized cone resistance  $Q_{tn}$  and the normalized sleeve friction  $F_r$  (Robertson 2009), together with estimates of the friction angle  $\varphi'$  in each soil unit (Kulhawy & Mayne 1990 in coarse grained soil, Mayne & Campanella 2005 in fine grained soils) and of the undrained shear strength  $c_u$  in clays (Lunne et al. 1997).

Fig. 5 shows the results obtained from SDMT C, carried out from the bank crest in area C, in terms of profiles with depth of various parameters provided by usual DMT interpretation (Marchetti 1980, Marchetti et al. 2001), i.e. the material index  $I_D$  (indicating soil type), the horizontal stress index  $K_D$  (related to stress history/*OCR*), the constrained modulus *M*, the undrained shear strength  $c_u$  (in clay), the friction angle  $\varphi'$  (in sand), as well as the profiles of the measured shear wave velocity  $V_S$  and the small strain shear modulus  $G_0$ , obtained as  $G_0 = \rho V_S^2$ .

Both CPTU and SDMT profiles (Figs 4 and 5) denote rather poor mechanical properties of the soils in the upper  $\approx 12$  m below the crest of the embankment (Units AR, B and C). In particular the sandysilty sediments of Unit B are characterized by low values of the horizontal stress index ( $K_D \approx 1-2$ ), which imply a low relative density  $D_R$ . Only the topmost 2-3 m of the embankment (Unit AR) show higher  $K_D$  values, presumably due to overconsolidation caused by desiccation-wetting cycles. The sands of Unit A, apart from sporadic thin layers having lower  $K_D$ , generally exhibit  $K_D \approx 3-5$ , thus denoting a medium relative density ( $D_R \approx 60\%$  according to Reyna & Chameau 1991). The interpretation of DMT results in the clay layers, excluding the shallow "crusts", indicates that the deposit is normally consolidated or slightly overconsolidated. The coefficient of earth pressure at rest in the fine-grained layers is generally  $K_0 \approx 0.6-0.7$ .

 $V_S$  measured by SDMT (Fig. 5) increases gradually with depth from  $\approx 150\text{-}200$  m/s in the topmost soil layers to  $\approx 300$  m/s at about 35 m depth. The profiles of the constrained modulus M (Marchetti 1980) indicate high compressibility of Units AR, B, C ( $M \approx 5$ -10 MPa) as well as of the deep clay layers, while the sands of Unit A are significantly less compressible. Differently from M, which refers to a "working strain" level (Marchetti et al. 2008), the values of the small strain shear modulus  $G_0$ , obtained from  $V_S$ measured in the same SDMT sounding, gradually increase with depth, without sharp contrasts between different soil layers.

The interpolation of the  $p_2$  values measured by SDMT indicated the presence of two distinct groundwater levels, thus confirming measurements



Figure 3. Stratigraphic model along the cross-section c-c', area C (Tonni et al. 2015a), including: borehole log; profiles of the corrected tip resistance  $q_i$  and the pore pressure u measured by CPTUs; profiles of the horizontal stress index  $K_D$  and the shear wave velocity  $V_S$  measured by SDMT.



Figure 4. Interpretation of results of CPTU 6 (area C).

in open standpipe piezometers and existing wells: indeed, the upper level is located in the sandy-silty sediments of the embankment core (Unit AR) and the underlying Unit B, generally at 4-5 m depth from the crest, whilst a lower piezometric level, governing pore pressures in the confined sandy layer of Unit A (the so-called "Acquifero Padano") can be identified at about 7-8 m depth from the crest.

#### **3 LIQUEFACTION ANALYSES**

#### 3.1 Procedure and seismic input data

Liquefaction analyses were carried out in each investigated area, in order to identify possible mechanisms responsible of the deformations and fractures observed on the crest of the embankment after the May 20, 2012 earthquake. The analyses were executed using a simplified dynamic approach, based on the comparison, at any depth, of the seismic demand on a soil layer generated by the earthquake (cyclic stress ratio *CSR*) and the capacity of the soil to resist liquefaction (cyclic resistance ratio *CRR*). When *CSR* is greater than *CRR* liquefaction may occur.



Figure 5. Interpretation of results of SDMT C (area C).

CSR was determined by 1-D ground seismic response analyses carried out using the code EERA (Bardet et al. 2000) in terms of total stresses, without taking into account the excess pore pressure build up typical of the liquefaction phenomenon. Details on the input data and results, obtained as part of the Working Group activity, can be found in Tonni et al. (2015a). The earthquake assumed as possible trigger of liquefaction was the May 20, 2012 main shock, recorded at 04:03 (local time), having local magnitude  $M_L = 5.9$  and epicentral distance  $R_{epi} = 7.5$  km from the Scortichino site. The main shock was followed, in about four minutes, by three aftershocks of  $M_L = 4.8, 4.8$  and 5.0 respectively and by nine shocks having  $M_L > 4$  within one hour. Since no ground motion recordings of this event were available in the area of Scortichino, the ground response analyses were carried out using four input accelerograms selected from the Italian earthquake database (ITACA 2011) by use of various search criteria (station on bedrock, moment magnitude  $M_w = 5.5-6.5$ ,  $R_{epi} = 5-10$  km). In addition, a near-fault accelerogram obtained for the April 6, 2009 L'Aquila 2009 earthquake was also considered. All the input accelerograms were scaled to a peak ground acceleration PGA = 0.183 g, estimated using an attenuation law (Bindi et al. 2011).

At each depth *CSR* was evaluated as:

$$CSR = \frac{\tau_{av}}{\sigma'_{v0}} = \frac{0.65\tau_{\max}}{\sigma'_{v0}}$$
(1)

where  $\tau_{max}$  is the maximum shear stress calculated by ground seismic response analysis (average of  $\tau_{max}$ calculated using different accelerograms),  $\tau_{av} = 0.65$  $\tau_{max}$  is the amplitude of the shear stress of the equivalent regular sequence, and  $\sigma'_{v0}$  is the effective overburden stress at the given depth.

*CSR* was then compared with the cyclic resistance ratio *CRR* estimated by use of various methods based on the SDMT parameters  $V_S$  and  $K_D$  (Tonni et al. 2015b), on the cone penetration resistance  $q_t$  from CPTU and from laboratory cyclic simple shear tests (Tonni et al. 2015a). The liquefaction safety factor *FSliq* at each depth was calculated as:

$$FS_{liq} = \frac{CRR}{CSR} = \frac{CRR_{M=7.5} \cdot MSF}{CSR}$$
(2)

where  $CRR_{M=7.5}$  is the cyclic resistance ratio for a reference magnitude  $M_w = 7.5$  (conventionally adopted in the simplified procedure) and *MSF* is a magnitude scaling factor. The analysis was carried out considering  $M_w = 6.14$ , equal to the maximum magnitude expected for a return period of 475 years in the seismogenetic zone in which Scortichino is located and similar to the magnitude of the May 20, 2012 main shock.

The "integral" liquefaction susceptibility at each test location was evaluated by means of the liquefaction potential index  $I_L$  (Iwasaki et al. 1982):

$$I_L = \int_{z=0}^{z_{crit}=20m} F(z) \cdot w(z) dz$$
(3)

where w(z) is a depth weighting factor and the function F(z) depends on the safety factor, according to Sonmez (2003).

#### 3.2 Evaluation of CRR from K<sub>D</sub> (SDMT)

In the last decades various *CRR-K<sub>D</sub>* correlations have been developed, including the most recent shown in Fig. 6, which appear to converge towards a narrow central band. Results of liquefaction analyses based on  $K_D$ , with *CRR<sub>M=7.5</sub>* estimated according to Monaco et al. (2005), Tsai et al. (2009) and Robertson (2012), were presented by Tonni et al. (2015b).

#### 3.3 Evaluation of CRR from $q_t$ (CPTU)

Results of liquefaction analyses based on CPTU, with  $CRR_{M=7.5}$  estimated from  $q_t$  according to Idriss & Boulanger (2004, 2006), were presented by Tonni et al. (2015a).

#### 3.4 Evaluation of CRR from the combination of $K_D$ (SDMT) & $q_t$ (CPTU)

As noted by Marchetti (2016), much of the interest on the CRR- $K_D$  correlation derives from the fact that the stress history increases significantly CRR and  $K_D$ , but only slightly the normalized cone resistance  $Q_{cn}$ . Hence it is possible that a correlation  $K_D$ -CRRwill be stricter than  $Q_{cn}$ -CRR.

The  $CRR-K_D$  correlation recommended by Marchetti (2016), identified by the label "RIB" in Fig. 6, was defined by combining the Idriss & Boulanger (2004, 2006)  $CRR-Q_{cn}$  correlation (Eq. 4a):

$$CRR = \exp \left[ (Q_{cn}/540) + (Q_{cn}/67)^2 - (Q_{cn}/80)^3 + (Q_{cn}/114)^4 - 3 \right]$$
(4a)

and the Robertson (2012) average  $Q_{cn} - K_D$  interrelationship (Eq. 4b):

$$Q_{cn} = 25 K_D \tag{4b}$$

A combined correlation for estimating *CRR* based at the same time on  $Q_{cn}$  and  $K_D$  (Eq. 5), plotted in the chart in Fig. 7 in the form  $CRR = f(Q_{cn}, K_D)$ , was obtained by Marchetti (2016) adopting as *CRR* the geometric average between a first *CRR* estimate obtained from  $Q_{cn}$  (Eq. 4a) and a second *CRR* estimate obtained from  $K_D$  (Eqs. 4a and 4b), namely:

Average  $CRR = [(CRR \text{ from } Q_{cn}) \cdot (CRR \text{ from } K_D)]^{0.5}$  (5)

#### 3.5 *Results and comments*

The results of liquefaction analyses based on different tests at the same location (area C) are compared in Fig. 8. The results provided by different methods



Figure 6. Recent *CRR-K<sub>D</sub>* correlations (Marchetti 2016).



Figure 7. Correlation for estimating *CRR* based on both  $Q_{cn}$  and  $K_D$ , for clean uncemented sand (Marchetti 2016).

based on the DMT parameter  $K_D$  (Monaco et al. 2005, Tsai et al. 2009, Robertson 2012, Marchetti 2016 "RIB"), denoted by white symbols in Fig. 8, indicate possible occurrence of liquefaction ( $FS_{liq} < 1$ ) in the silty-sandy soils at the base of the embankment (Unit B), at local depths from the crest between about 5 to 9 m, while no significant liquefaction is detected in the deeper sands (Unit A). The liquefaction potential index  $I_L$  estimated from  $K_D$  by all methods is "high".

The results of the analyses based on  $q_t$  from CPTU (Idriss & Boulanger 2004, 2006), denoted by black symbols in Fig. 8, signal the presence of a liquefiable layer, having much lower thickness than indicated by  $K_D$ , within the silty sand of Unit B. Differently from  $K_D$ , the analysis based on  $q_t$  suggests generalized liquefaction in the deeper sands (Unit A). The liquefaction potential index  $I_L$  is "moderate" to "high", i.e. lower than indicated by  $K_D$ .

The results obtained by the method proposed by Marchetti (2016), based on both  $K_D$  from SDMT and

 $q_t$  from CPTU (grey symbols in Fig. 8), indicate the presence of a liquefiable layer of lower thickness than indicated by  $K_D$  alone within Unit B, in agreement with the analysis based on  $q_t$  alone. At the same time, this method tends to exclude significant liquefaction in Unit A, in agreement with the analyses based on  $K_D$  alone. The liquefaction potential index  $I_L$  is "low", i.e. substantially lower than indicated by methods based on  $K_D$  and  $q_t$  alone.

To note that all the *CRR* correlations based on  $K_D$  (Monaco et al. 2005, Tsai et al. 2009, Robertson 2012, Marchetti 2016) are valid for clean sand, without any correction for fines content. Hence the *CRR* estimated from  $K_D$  in the sandy-silty layers (Units AR and B) are probably somewhat underestimated (though the low plasticity of fines in these layers should not involve a substantial increase in *CRR*), while in the clean sands of Unit A the *CRR* estimated from  $K_D$  are presumably realistic.

In Fig. 8 the results of the analyses based on  $K_D$  and  $q_t$  – alone and combined – are compared with the results obtained by Tonni et al. (2015b) using the correlations based on  $V_S$  by Andrus & Stokoe (2000) and Kayen et al. (2013). The analyses based on  $V_S$  generally indicate minor liquefaction ("low"  $I_L$ ).



Figure 8. Area C. Results of liquefaction analyses based on the horizontal stress index  $K_D$  (SDMT), on the cone penetration resistance  $q_t$  (CPTU) and on the combination  $K_D$  (SDMT) &  $q_t$  (CPTU), compared with results obtained by methods based on the shear wave velocity  $V_S$  and by laboratory cyclic simple shear tests (CSS).

Fig. 8 also shows the value of  $FS_{liq}$  obtained by a laboratory cyclic simple shear test (CSS) performed on a silty-sandy sample taken in borehole S5 at 6.00-6.60 m depth. This result ( $FS_{liq} = 1$ ) confirms the possible occurrence of liquefaction in the silty-sandy layer.

### 4 CONCLUSIONS

The results of liquefaction analyses carried out using simplified methods based on the DMT horizontal stress index  $K_D$ , in agreement with well-established methods based on the CPT cone penetration resistance  $q_t$ , suggest that plausibly local liquefaction phenomena, of variable extent, may have been induced by the May 20, 2012 earthquake in the sandysilty soils below the Scortichino canal levee.

In the case illustrated in the paper, the use of a combined correlation for estimating *CRR* based at the same time on CPT- $q_t$  and DMT- $K_D$  (Marchetti 2016) has confirmed the probable occurrence of liquefaction. However the estimated overall liquefaction susceptibility, represented by the liquefaction potential index  $I_L$ , is lower than indicated by methods based on both  $K_D$  alone and  $q_t$  alone. This result is in reasonable agreement with field observations.

As noted by Marchetti (2016), it is expectable that an estimate based at the same time on two measured parameters is more accurate than estimates based on just one parameter, and incorporating the DMT stress history parameter  $K_D$  into the liquefaction correlations should possibly reduce the uncertainty in estimating *CRR*. Considerable additional research is obviously necessary, especially if the sand is not clean, uncemented sand.

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